

Abstract

To confirm the usefulness of the chemistry of detrital heavy minerals, the provenance of two collisional areas were studied, and the paleogeography before collision and the uplift history of the collisional zone were reconstructed mainly based on the chemistry of detrital heavy minerals. The first is the collision between the Honshu and Izu arcs since Middle Miocene, and the paleogeography before the collision was reconstructed. The second is the collision between the Eurasia Plate and Indian Subcontinent since the earliest Eocene, and the uplift history of the Himalayas was reconstructed.

The latest Oligocene to early Middle Miocene strata in the Mineoka area and the late Eocene to early Middle Miocene strata in the Setogawa area were investigated to reconstruct the paleogeography around the Izu arc collisional zone. The sandstones of the Eocene to Oligocene Mikura Group in the Setogawa area were plotted in the field of "Recycled Orogen" or "Dissected Arc" provenance and they yield the detrital chromian spinels and garnets whose chemistries are similar to those of the spinels in serpentinites of the Kurosegawa Belt and garnets in metamorphic rocks and granite in the Ryoke Belt, respectively. The detritus of the Mineoka, Setogawa and Kurami Group had a similar provenance to that of the Mikura Group except for the origin of detrital chromian spinels. From these strata, detrital chromian spinels whose chemistries are similar to those in the Circum-Izu Massif serpentinites were obtained. Most of the Miocene strata in the Mineoka-Setogawa areas yield detrital chromian spinels whose chemistries are similar to those in the Mineoka type serpentinites, while only the conglomerates in the Mineoka and Setogawa Groups contain detrital chromian spinels whose chemistries are similar to those in the Hayama type serpentinites. It is inferred that the Hayama type serpentinites of the Circum-Izu Massif serpentinites was exposed at first and then the emplacement of the Mineoka type serpentinites occurred. The origin of the Circum-Izu Massif

serpentinites was considered to be backarc peridotites, and the existence of an arc to the east of the Mineoka-Setogawa areas in the Early Miocene was also inferred.

The metamorphic rocks and granites in the Ryoke Belt had supplied detrital garnet to the early Early Miocene to early Middle Miocene strata in the Mineoka and Setogawa area except for the Hota Group. The Hota Group sandstone can be divided into two groups. One group is distributed to the south of the Mineoka Mountains at present, and their detritus had been mostly derived from the volcanoclastic rocks. The chemistry of detrital clinopyroxenes obtained from this group sandstones show the orogenic calcalkaline rock origin. The volcanic activity which supplied detritus to the Hota Group correspond to the northeast Honshu arc or the inferred oceanic island arc. The other group is distributed to the north of the Mineoka Mountains at present. Their sandstones are quartz-feldspathic, and contain a large amount of detrital garnets whose chemistry is similar to those in the continental basement. Thus, there is a possibility that the detritus of the second group was supplied from the continental basement directly.

The modal composition of sandstone and the assemblage of the detrital heavy minerals change between the sandstone from the lower and upper parts of the upper Kushigatayama Subgroup of the early Middle Miocene Koma Group. The chemistries of detrital clinopyroxenes are similar to those in the volcanic rocks in the Izu arc. It is inferred that the change of the provenance observed in the Koma Group was caused by the collision of the Honshu and Izu arc.

The strata from the late Eocene to the present in the Bengal basin were also investigated to reconstruct the uplift history of the Himalayas. From the late Eocene Kopili Formation in Bangladesh, detrital chromian spinels whose chemistries are similar to those of the spinels in the Yarlung-Zangbo ophiolite and intra-plate basalts such as those in the Deccan Traps were obtained. Because the Middle Eocene Sylhet Limestone in Bangladesh contains no detritus, it is inferred that the first supply of detritus from the Himalayas to the Bengal Basin occurred in the late Eocene, and a small amount of the detritus from the Indian Subcontinent had been also supplied in

this period. The Yarlung-Zangbo ophiolite have supplied detritus until the present, however, the detrital chromian spinels whose chemistries are similar to the intra-plate basalts are rarely included in the strata younger than the Kopili Formation.

The detrital garnets were also found from all the strata in the Bengal basin investigated in this study. The chemistry of detrital garnets had changed two times from the Eocene to the present. The late Eocene Kopili and the Oligocene Barail Formations sandstones contain detrital garnets whose chemistries are similar to those in the Transhimalaya and Higher and Tethys Himalayas. The first change occurred between the Oligocene Barail Formation and the Early Miocene Bhuban Formation. The pyrope-rich almandine was predominant in the detrital garnets from the Barail Formation, while the both the pyrope-rich almandine and spessartine-rich almandine was predominant in the detrital garnets from the Bhuban Formation. This change seems to be caused by active dissection of the Tertiary granitic rocks in the Himalayas. The second change occurred between the Bhuban Formation and the overlying Boka Bil Formation. The Boka Bil Formation sandstones contain a large amount of especially spessartine-rich almandine and grandite. It is probable that this change was caused by the start of dissection of the Lesser Himalayas. These characteristics are also common in the sandstones and sands from late Miocene to the present.

Key words

sandstone, shale, detrital chromian spinel, detrital garnet, detrital clinopyroxene, collisional zone, Mineoka-Setogawa area, Himalayas-Bengal system

List of Tables

Table 1	Modal composition of sandstones in the Mineoka area based on traditional method	53
Table 2	Modal composition of sandstones in the Mineoka area based on Gazzi-Dickinson method	54
Table 3	Clast composition of the conglomerate of the Haccho Formation at Fusada	58
Table 4	Modal composition of sandstones in the Setogawa area based on traditional method	61
Table 5	Modal composition of sandstones in the Setogawa area based on Gazzi-Dickinson method	62
Table 6	Frequency of occurrence of detrital chromian spinels in the Mineoka area	67
Table 7	Frequency of occurrence of detrital chromian spinels in the Setogawa area	69
Table 8	Modal composition of sandstones and sands in the Bengal basin based on traditional method	149
Table 9	Modal composition of sandstones and sands in the Bengal basin based on Gazzi-Dickinson method	150

List of Figures

Fig. 1	Geologic map around the Izu collisional zone	3
Fig. 2	Geologic map around the Himalayas and sites of piston core sampling points	4
Fig. 3	Location map of the Mineoka and Setogawa areas	5
Fig. 4	Arrangement of equipment for heavy liquid separation	10
Fig. 5	Historical review of the stratigraphy of the southern part of the Boso Peninsula	17
Fig. 6	Geological map around the Mineoka Mountains in the southern part of the Boso Peninsula	25
Fig. 7	Locations of route map in the Mineoka area	28
Fig. 8	Route map of Motona area	29
Fig. 9	Route map of the Yoshinodai River, branch of the Hiratsuka River, in the Fusada area	30
Fig. 10	Correlation diagram of the groups in the Setogawa area	34
Fig. 11	Simplified geological map of the Setogawa area	35
Fig. 12	Geologic map of the southern distributional area of the Mikura Group	36
Fig. 13	Geologic map of the Setogawa Group in the middle reaches of the Abe River	39
Fig. 14	Geologic map of the Kurami Group	43
Fig. 15	Geologic map of the Koma Mountains	45
Fig. 16	Route map along the Yohka Beach	48
Fig. 17	Locations of tectonic blocks at Mineokasengen	49
Fig. 18	Sketch map of tectonic blocks along the Ara River at Hegurinaka	50
Fig. 19	Q-F-R diagrams of the sandstone in the Mineoka area	55
Fig. 20	Q-F-L and Q _m -F-L _t diagrams of the sandstones in the Mineoka area based on the Gazzi-Dickinson method	56

Fig. 21	Clast composition of the conglomerate of the Haccho Formation at Fusada	59
Fig. 22	Q·F·R diagrams of the sandstones in the Setogawa area	63
Fig. 23	Q·F·L and Q _m ·F·L _t diagrams of the sandstones in the Setogawa area based on the Gazzi-Dickinson method	64
Fig. 24	Chemical compositions of detrital chromian spinels from sandstones of the Haccho Formation of the Mineoka Group	71
Fig. 25	Chemical compositions of detrital chromian spinels from conglomerates of the Haccho Formation of the Mineoka Group	72
Fig. 26	Chemical compositions of detrital chromian spinels from the Hota Group	73
Fig. 27	Chemical compositions of detrital chromian spinels from the Sakuma Group	74
Fig. 28	Chemical compositions of detrital chromian spinels from the serpentine sandstone	75
Fig. 29	Chemical compositions of detrital chromian spinels from the leucocratic sandstone	76
Fig. 30	Chemical compositions of detrital chromian spinels from the Mikura Group	77
Fig. 31	Chemical compositions of detrital chromian spinels from the Setogawa Group	78
Fig. 32	Chemical compositions of detrital chromian spinels from the Kurami Group	79
Fig. 33	Chemical compositions of detrital chromian spinels from the Koma Group	80
Fig. 34	Chemical compositions of detrital garnets from the Haccho Formation of the Mineoka Group	86
Fig. 35	Chemical compositions of detrital garnets from the Hota Group	87
Fig. 36	Chemical compositions of detrital garnets from the Sakuma Group	88

Fig. 37	Chemical compositions of detrital garnets from the Futatsuyama Formation	89
Fig. 38	Chemical compositions of detrital garnets from the leucocratic sandstone	90
Fig. 39	Chemical compositions of detrital garnets from the Mikura Group	91
Fig. 40	Chemical compositions of detrital garnets from the Setogawa Group	92
Fig. 41	Chemical compositions of detrital garnets from the Kurami Group	93
Fig. 42	Chemical compositions of detrital garnets from the Koma Group	94
Fig. 43	Di-En-Hd-Fs diagram for detrital clinopyroxenes from the Mineoka area	99
Fig. 44	Di-En-Hd-Fs diagram for detrital clinopyroxenes from the Mineoka and Setogawa areas	100
Fig. 45	Discrimination diagrams for detrital clinopyroxenes from conglomerate of the Haccho Formation, Mineoka Group	118
Fig. 46	Discrimination diagrams for detrital clinopyroxenes from the Hota Group	119
Fig. 47	Discrimination diagrams for detrital clinopyroxenes from the Sakuma Group	120
Fig. 48	Discrimination diagrams for detrital clinopyroxenes from the Futatsuyama Formation	121
Fig. 49	Discrimination diagrams for detrital clinopyroxenes from the serpentine sandstone at Yohka Beach	122
Fig. 50	Discrimination diagrams for detrital clinopyroxenes from the serpentine sandstone at Mineokasengen	123
Fig. 51	Discrimination diagrams for detrital clinopyroxenes from the Kushigatayama Subgroup of the Koma Group	125
Fig. 52	Reconstructed paleogeography in the late Eocene to early Oligocene time for the Mineoka-Setogawa area	128
Fig. 53	Reconstructed paleogeography in the earliest early Miocene time	

	for the Mineoka-Setogawa area	129
Fig. 54	Reconstructed paleogeography in the early Early Miocene time for the Mineoka-Setogawa area	130
Fig. 55	Reconstructed paleogeography in the late Early Miocene time for the Mineoka-Setogawa area	131
Fig. 56	Reconstructed paleogeography in the early Middle Miocene time for the Mineoka-Setogawa area	132
Fig. 57	Geologic map of Bangladesh	141
Fig. 58	Stratigraphy of the Sylhet and Chittagong areas	142
Fig. 59	Geological map of the northern part of the Sylhet area	145
Fig. 60	Geologic map of the Chittagong area	146
Fig. 61	Lithology of piston cores from the Bengal Fan	147
Fig. 62	Q-F-R diagrams of the sandstones and sands in the Bengal basin	151
Fig. 63	Q-F-L and Qm-F-Lt diagrams of the sandstones and sands in the Bengal basin based on the Gazzi-Dickinson method	152
Fig. 64	Chemical compositions of detrital chromian spinels from the Kopili Formation	155
Fig. 65	Chemical compositions of detrital chromian spinels from the Barail Formation	156
Fig. 66	Chemical compositions of detrital chromian spinels from the Bhuban Formation	157
Fig. 67	Chemical compositions of detrital chromian spinels from the Boka Bil Formation	158
Fig. 68	Chemical compositions of detrital chromian spinels from the Tipam Group	159
Fig. 69	Chemical compositions of detrital chromian spinels from the Dupi Tila Formation	160
Fig. 70	Chemical compositions of detrital chromian spinels from beach sand in the Chittagong City	161

Fig. 71	Chemical compositions of detrital chromian spinels from piston core sample of BP1	162
Fig. 72	Chemical compositions of detrital chromian spinels from piston core sample of BP3	163
Fig. 73	Chemical compositions of detrital chromian spinels from piston core sample of BP4	164
Fig. 74	Chemical compositions of detrital garnets from the Kopili Formation	168
Fig. 75	Chemical compositions of detrital garnets from the Barail Formation	169
Fig. 76	Chemical compositions of detrital garnets from the Bhuban Formation	170
Fig. 77	Chemical compositions of detrital garnets from the Boka Bil Formation	171
Fig. 78	Chemical compositions of detrital garnets from the Tipam Group	172
Fig. 79	Chemical compositions of detrital garnets from the Dupi Tila Formation	173
Fig. 80	Chemical compositions of detrital garnets from beach sand in the Chittagong City	174
Fig. 81	Chemical compositions of detrital garnets from piston core sample of BP1	175
Fig. 82	Chemical compositions of detrital garnets from piston core sample of BP3	176
Fig. 83	Chemical compositions of detrital garnets from piston core sample of BP4	177
Fig. 84	Model of the provenance changes of sediments in the Bengal basin	187